Review

Impact of land use change on greenhouse gases emissions in peatland: a review

Ryusuke Hatano

Research Faculty of Agriculture, Hokkaido University, Kita 9, Nishi 9, Kita-ku, Sapporo 060-8589, Japan

Received September 10, 2018; accepted December 22, 2018

Abstract. Peatland is a significant storage of carbon and nitrogen on the earth's surface. This paper reviews the impacts of changes in water table level and mineral nitrogen associated with human activities on greenhouse gases emissions in tropical peatland and northern boreal and temperate peatland, and evaluates the optimal water table level to minimize greenhouse gases emissions. CH₄ emission increased significantly with the rise of ground water table level above -20 cm, and larger in northern peatland with plant mediated CH₄ emission than tropical peatland with plant mediated oxygen supply. However, forest disturbance by fire in tropical peatland increased CH₄ flux to the similar level in northern peatlands (8.3 mg C m⁻² h⁻¹) due to stagnant surface water associated with the peat subsidence. On the other hand, CO₂ and N₂O emissions were significantly lager in tropical peatland than in northern peatland especially due to nitrogen fertilization. CO2 and N2O emissions increased with falling ground water table level below -40 to -80 cm (19 Mg C ha⁻¹ y⁻¹ for CO_2 and 700 kg N ha⁻¹ y⁻¹ for N₂O). Total global warming potential was significantly low in the ground water table level from -20 and -40 cm.

K eywords: carbon dioxide, ground water table, methane, nitrogen fertilizer, nitrous oxide, peatland

INTRODUCTION

Peatland is generated in wetland due to lower microbial decomposition than plant production. It has been developed at the organic matter accumulation rate of 100 to 200 kg C ha⁻¹ y⁻¹ for past 10 000 years, and the area is 4.4 million km², 4 million km² in northern boreal and temperate region and 0.4 million km² in tropical region (Yu *et al.*, 2010). Although the area occupies only 3% of terrestrial area, it contains 612 Gt C (562 Gt C in northern boreal and temperate peatland, 50 Gt C in tropical peatland), accounting for 25% of total soil carbon. However, peatland is different

among regions. Northern boreal and temperate peatland is usually covered by small vegetation, sphagnum (bog) and carex and shrub (fen). On the other hand, tropical peatland is covered by trees (peat swamp forest). Northern boreal and temperate peat profile is composed of fine fibric vegetation materials, it makes fine pore system. On the other hand, tropical peat profile contains woody materials, it makes macropore system. Therefore, natural tropical peatland has high water permeability (Melling, 2016).

As peatland is composed of organic matter, peat decomposition produces a significant amount of greenhouse gases (GHGs); carbon dioxide (CO_2) aerobically, methane (CH_4) anaerobically and nitrous oxide (N₂O) both aerobically and anaerobically. Global warming potential (GWP) of CH4 and N₂O is 28 and 265 times larger than CO₂, respectively, over a 100-year period, and those GWP values do not include climate-carbon feedbacks (IPCC, 2014), and N₂O is an important ozone-depleting substance emitted into the atmosphere (UNEP, 2013). Land use change due to disturbance of peatland by fire and clearance of vegetation and the development of peatland for agricultural land use by drainage, plowing and fertilization influence GHGs emissions (Page and Baird, 2016; van der Werf et al., 2009). Falling ground water level increases CO₂ (Couwenberg et al., 2010; Ishikura et al., 2017, 2018), but decreases CH₄ emission (Couwenberg et al., 2010; Ishikura et al., 2018). CO_2 emission increases with the increase of temperature (Lloyd and Taylor, 1994). On the other hand, CH₄ emission is controlled by redox potential in soil (Takai, 1970). Therefore, CH₄ emission is influenced by the ground water level of peatland, the water permeability of peat soil. Lower

Int. Agrophys., 2019, 33, 167-173 doi: 10.31545/intagr/109238

^{*}Corresponding author e-mail: hatano@chem.agr.hokudai.ac.jp

permeable peat with finer pore system may produce higher CH_4 emission. Concerning N₂O emission, it increases with increase of organic matter decomposition and nitrogen fertilization application (Mu *et al.*, 2009). Although it is not clearly mentioned the relationship between N₂O emission and ground water table level, nitrogen mineralization with falling ground water level may increase N₂O emission. However, nitrogen fertilizer application inhibits microbial organic matter decomposition, although it stimulates plant root respiration (Zhou *et al.*, 2014).

These mentioned above suggest that CH_4 , and N_2O emissions together with CO_2 emissions can be controlled by the ground water table level. The aim of this paper is to discuss about optimal ground water table level minimizing CO_2 , N_2O and CH_4 emissions by reviewing the published papers.

Wetland is a major source of CH₄. Natural wetland emits 217 Tg CH₄ y⁻¹, accounting for 63% of total natural CH₄ emission, and rice paddy field emits 36 Tg CH₄ y⁻¹ accounting for 11% of total anthropogenic CH4 emission (IPCC, 2013). CH₄ is produced in reductive subsoil, transported to top soil by molecular diffusion or ebullition as gas bubbles, and emitted from the surface of the soil or the plant leaves through aerenchyma from the roots (Schütz et al., 1989), but 90% of the CH₄ produced in soil is oxidized during the transportation in the rhizosphere, either by oxygen released from plant roots or by other electron acceptors such as Fe(III) and SO42- (Kolb and Horn, 2012). Ground water table level is major controlling factor of CH4 emission (Couwenberg et al., 2010; Ishikura et al., 2018). CH₄ emission increases significantly when ground water table level rise above -20 cm, however, CH₄ emission was lower in tropical peatland than in northern boreal and temperate peatland (Couwenberg et al., 2010) (Fig. 1). This is probably due to rapid fluctuation of ground water table level associated with better water permeability in tropical peatland (Takahashi, 1999). CH₄ is produced in anaerobic condition by methanogenic bacteria using acetate and $CO_2 + H_2$ after the sequential utilization of a series of electron acceptors of oxygen, NO₃, Mn(IV), Fe(III), and SO₄²⁻ (Takai, 1970). Therefore, peat soil with lower permeability, lower fertility and lower contamination of mineral soil produces CH₄ more easily due to less contain of electron acceptors. Recent global model uses 0.2 to 0.25 for the CH_4/CO_2 mole ratio in boreal natural peatland and 0.0052 in tropical natural peatland (Spahni et al., 2011). This is because of strong oxidizing power in natural tropical peatland induced by the plant mediate oxygen supply through aerial roots (Adji et al., 2014) and fast water flow in well water permeable peat layer (Kelly et al., 2014). Peat fire or clear tree cutting increases CH₄ emission in tropical peatland due to the disappearance of plant mediate oxygen supply and the rise of ground water table level with peat subsidence, while drainage decreases CH₄ emission due to the fall of ground water table level (Adji et al., 2014; Jauhiainen et al., 2008)



Fig. 1. Relationship between CH_4 flux and ground water table level. Dashed line shows -20 cm of ground water table level at which CH_4 flux clearly increased (modified from Couwenberg *et al.*, 2010).



Fig. 2. Relationship between CH_4 flux and ground water table level in a tropical peatland (modified from Adji *et al.*, 2014).

(Fig. 2). Maximum CH₄ flux was 8.3 mg C m⁻² h⁻¹ in burnt area. This is significantly large value within tropical peatland other than rice paddy fields (2.6-11 (Hadi *et al.*, 2005) and up to 26 mg C m⁻² h⁻¹ (Furukawa *et al.*, 2005)) and is similar to the maximum of CH₄ emission in boreal and temperate peatland (Couwenberg *et al.*, 2010).

CH₄ flux from tree stem in natural tropical peatland is reported, but the value was only 0.021 mg C m⁻² h⁻¹ (Pangala *et al.*, 2013). On the other hand, in northern boreal and temperate peatlands, pond vegetation accelerates CH₄ emission due to transportation through the aerenchyma. An observation in a thermokarst depression of 63.7 ha including pond, wet grassland, dry grassland near forest in Yakutia for four years from 2006 to 2009 showed that 96.3 ± 5.4% of total CH₄ emission was emitted from thermokarst pond (the area was 60.6 ± 19.2% of the total area) and plant mediated CH₄ emission from the pond vegetation accounted for 57.9 ± 23.2% of total CH₄ emission (Desyatkin *et al.*, 2014).

Natural wetland is not a source of N_2O , rather a sink of N_2O . This is because of slow N mineralization, low nitrification activity, and high denitrification activity in

the anaerobic condition (Kolb and Horn, 2012). But, once peatland is drained and agriculturally used, the peatland emits N₂O significantly. Significantly large amount of N₂O emission more than 200 kg N ha⁻¹ was recorded in NO₃-N accumulated upland fields in tropical peatland especially in wet season (Takakai et al., 2006; Toma et al., 2011) and 9 to 14 kg N ha⁻¹ was found even in tundra peatland (Repo et al., 2009). This is because of nitrogen application for crop production and incomplete denitrification with NO₃-N leaching (Kolb and Horn, 2012). N₂O emission from agricultural soil accounts for 34% of total anthropogenic N₂O emission of 6.9 Tg N y⁻¹, although N₂O emission from natural ecosystem is estimated to be 11.0 TgN y⁻¹ (IPCC, 2013). Denitrification is the major N₂O production process in most soils, although N₂O is produced in nitrification process (Šimek et al., 2002). Denitrification rate in soil is increased with the increases of NO₃⁻ content, available organic carbon content, temperature, soil moisture (Saggar et al., 2013), and decrease of soil pH (Mukumbuta et al., 2018). N₂O is produced as an intermittent product through the process of NO_3^- reduction to N_2 by nitrate respiration of denitrifers. Therefore, presence of NO_3^- and available organic carbon is indispensable for N₂O production in denitrification. N₂O is mainly produced in the range of water-filled pore space (WFPS) between 60 to 100%, because optimal WFPS for nitrification producing NO₃⁻ is around 60 while 100% of WFPS (water saturation) is optimal for denitirification (Linn and Doran, 1984). Peat soil has high amount of pore space, and drained and plowed agricultural peat soil may have ideal WFPS for nitrification and denitrification processes. Soil pH is also important for denitrification. The N₂O reductase is less active at low pH compared to other enzyme activities in denitrification process (McMillan et al., 2016), which leads to increase of N₂O emission with increase of acidity in a Japanese Andosol (Mukumbuta et al., 2018). Peat soil is acidic due to rich in organic acids, which can increase of N₂O emission especially in cropland.

Mu *et al.* (2014) shows that N_2O emission increased with the increase of CO_2 emission by peat decomposition using the 122 published data of N₂O and CO₂ emissions measured at forests, grasslands and croplands in northern and tropical peatlands. The emission data were obtained by closed chamber methods. However, some of peat decomposition (Rh, kg C ha⁻¹ y⁻¹) was estimated from soil respiration (Rt, kg C ha⁻¹ y⁻¹) including root respiration using an empirical equation, Rh=10exp (0.22+0.87 ln(Rt/10)), proposed by Bond-Lamberty and Thomson (2010). It was shown that N₂O emission was significantly correlated with mineral N input (sum of application rate of nitrogen fertilizer and mineralized nitrogen estimated by dividing Rh by soil C/N ratio), ground water table level, and soil pH. However, data sets from tropical peatlands were only 12 from Malaysia and Indonesia. Therefore, newly published 30 data sets from tropical peatlands were added in this paper (Melling et al., 2007; Takakai et al., 2006; Toma et al., 2011), total 152 data sets including 87 data from unfertilized northern peatland, 23 data from fertilized northern peatland, 26 data from unfertilized tropical peatland and 16 data from fertilized tropical peatland were obtained. In this paper, comparison between tropical peatland and northern boreal and temperate peatland was conducted in terms of the relations between N₂O and CO₂ emissions and the effect of fertilization on the emissions. Statistical analyses were performed with Excel Statistics version 5.0 (Esumi, Tokyo, Japan). The differences in the emissions between fertilization and unfertilization and between tropical peatland and northern boreal and temperate peatland were analyzed with a two-way analysis of variance (ANOVA) and Tukey test.

CO₂ emission was significantly higher in tropical peatland than northern boreal and temperate peatland, and increased with N fertilization significantly (3815 ± 2900) and 4822 ± 2313 kg C ha⁻¹ y⁻¹ in unfertilized and fertilized northern boreal and temperate peatland, respectively, and 7382 ± 3558 and 13001 ± 3027 kg C ha⁻¹ y⁻¹ in unfertilized and fertilized tropical peatland, respectively) (Fig. 3). N₂O emission increased with N fertilization significantly, especially in tropical peatland $(4.93 \pm 10.04 \text{ and } 17.76 \pm 21.15$ kg N ha⁻¹ y⁻¹ in unfertilized and fertilized northern boreal and temperate peatland, respectively, and 8.08 ± 15.69 and 178.59 ± 218.28 kg N ha⁻¹ y⁻¹ in unfertilized and fertilized tropical peatland, respectively) (Fig. 4). Furthermore, in all peatlands, CO₂ and N₂O emissions increased with the fall of ground water table level, and much larger increase was found in nitrogen fertilized peatlands (Figs 5, 6). Increase of CO₂ emission with the fall of water table level was larger in tropical peatland than in northern boreal and temperate peatland, and maximum CO₂ emission of 19 Mg C ha⁻¹ y⁻¹ was found at -80 cm of water table level (Fig. 5). Increase of N2O emission with the fall of water table level was distinct below -40 cm of water table level, and especially in tropical peatland, maximum N₂O emission of 700 kg N ha⁻¹ y⁻¹ was found in -60 to -70 cm of water table level (Fig. 6).



Fig. 3. Comparison of annual CO_2 emission with peat decomposition in terms of nitrogen fertilization and climate zone. Error bar reveals standard deviation. Different letters reveals significant difference among the treatments of both fertilization and climate zone (data compiled by Mu *et al.*, 2014 and combined with those from Melling *et al.*, 2007; Takakai *et al.*, 2006; and Toma *et al.*, 2011).



Fig. 4. Comparison of annual N₂O emission in terms of nitrogen fertilization and climate zone. Error bar reveals standard deviation. Different character reveals significant difference among the treatments of both fertilization and climate zone (data compiled by Mu *et al.*, 2014 and combined with those from Melling *et al.*, 2007; Takakai *et al.*, 2006 and Toma *et al.*, 2011).



Fig. 5. Relationship between annual CO_2 emission and annual mean ground water table level in peatlands (data compiled by Mu *et al.*, 2014 and combined with those from Melling *et al.*, 2007; Takakai *et al.*, 2006 and Toma *et al.*, 2011).



Fig. 6. Relationship between annual N₂O emission and annual mean ground water table level in peatlands (data compiled by Mu *et al.*, 2014 and combined with those from Melling *et al.*, 2007; Takakai *et al.*, 2006 and Toma *et al.*, 2011).

There was a significant relationship between mineral nitrogen input and N_2O emission (Fig. 7). This indicates that increase of peat decomposition with a fall of water table level and N fertilizer application increases N_2O emission.

In an upland field of tropical peatland where large N_2O emission was observed, there was a significant relationship between NO₃-N content in top soil and N₂O flux at more than 60%, of WFPS suggesting N₂O emission was caused by denitrification after mineralization and nitrification (Takakai *et al.*, 2006). Soil microbes which adapted to low pH of tropical peatland and obtained nitrate respiration ability, fungi (*Fusarium oxysporum* and *Neocosmospora vasinfecta*) (Yanai *et al.*, 2007) and bacteria (*Janthinobacterium*) (Hashidoko *et al.*, 2008), were identified.



Fig. 7. Relationship between annual N₂O emission and annual mineral nitrogen input in peatlands (data compiled by Mu *et al.*, 2014 and combined with those from Melling *et al.*, 2007; Takakai *et al.*, 2006 and Toma *et al.*, 2011).

From the findings above, ground water table level is major controlling factor of CH₄ and N₂O emissions from peatlands. Ground water table level rising above -20 cm increases CH₄ emission. The increase of CH₄ emission is lager in northern boreal and temperate peatland than in tropical peatland due to higher water permeability in tropical peatland. However, loss of natural vegetation with aerial roots and peat subsidence by disturbance especially by fire in tropical peatland increases CH₄ emission significantly due to loss of plant mediated oxygen supply and decrease of redox potential under stagnant water on the soil surface. On the other hand, ground water level from -40 to -70 cm stimulates N_2O emission together with CO_2 emission. Nitrogen fertilizer application increases N2O and CO₂ emissions significantly especially in tropical peatland. Overall emission data show the lowest emissions of CH₄ and N₂O are achieved in the range of -20 to -40 cm of ground water level. Figure 8 shows the mean and standard deviation of total GWP of CO2, N2O and CH4 in three ranges of water table level, higher than -20 cm, from -20 to -40 cm and lower than -40 cm. The GWP (kg CO_2 ha⁻¹ y⁻¹)



Fig. 8. Comparison of annual global warming potential (GWP) in terms of ground water table level (WTL) and climate zone. Error bar reveals standard deviation of total GWP. Different character reveals significant difference among the treatments of ground water table level and climate zone.

of each gas was obtained from the data of Fig. 1 for CH₄ and Fig. 5 for CO_2 and Fig. 6 for N_2O_2 , using the factors reported by IPCC 2014, as $GWP_{CO_2} = CO_2$ emission (kg C ha⁻¹ y⁻¹)×44/12, GWP_{CH₄}=CH₄ emission (kg C ha⁻¹ y⁻¹) $\times 16/12 \times 28$ and GWP_{N20}=N₂O emission (kg N ha⁻¹ y⁻¹) ×44/28×265. Unit of CH₄ emission in Fig. 1 was converted to calculate GWP_{CH_4} . Northern boreal and temperate peatland showed the significantly lowest total GWP in the range of water table level from -20 to -40 cm, and tropical peatland showed that total GWP was significantly lower in the range of water table level from -20 to -40 cm than the water table level lower than -40 cm, and there was no significant difference of total GWP between the ranges of water table level higher than -20 cm and from -20 to -40 cm. However, total GWP was significantly higher in tropical peatland than in northern boreal and temperate peatland. In both tropical and northern boreal and temperate peatlands, contribution of CO₂ emission to total GWP was highest in the all ranges of water table level. CH₄ emission showed larger contribution in the higher range of water table level, and especially in northern boreal and temperate peatland, contribution of CH₄ emission was similar to that of CO₂ emission. On the other hand, contribution of N₂O emission was higher in the lower range of water table level, and especially in tropical peatland, contribution of N₂O emission was similar to that of CO₂ emission.

SUMMARY

Water table level from -20 cm to -40 cm was optimal to minimize CH_4 , CO_2 and N_2O emissions in peatlands. Tropical peatland was lower in CH_4 emission than northern boreal and temperate peatland, however, stagnant surface water in subsided and fired tropical peatland induced high CH_4 flux. On the other hand, tropical peatland was higher in N_2O emission than northern boreal and temperate peatland. Falling ground water table level stimulated organic matter decomposition significantly, leading to increase of mineralized N resulting in increase of N₂O emission. Application of nitrogen fertilizer especially in tropical peatland increases N₂O emission significantly due to significant increase of organic matter decomposition.

Conflict of interest: The Authors do not declare conflict of interest.

REFERENCES

- Adji F.F., Hamada Y., Darang U., Limin S.H., and Hatano R., 2014. Effect of plant-mediated oxygen supply and drainage on greenhouse gas emission from a tropical peatland in Central Kalimantan, Indonesia. Soil Sci. Plant Nutrition, 60, 216-230. https://doi.org/10.1080/00380768.2013.872019
- Bond-Lamberty B., and Thomson A., 2010. A global database of soil respiration data. Biogeosciences, 7, 1915-1926. https://doi.org/10.5194/bg-7-1915-2010
- Couwenberg J., Dommain R., and Joosten H., 2010. Greenhouse gas fluxes from tropical peatlands in south-east Asia. Global Change Biol., 16, 1715-1732. https://doi. org/10.1111/j.1365-2486.2009.02016.x
- Desyatkin A.R., Takakai F., and Hatano R., 2014. Flood effect on CH₄ emission from the alas in Central Yakutia, East Siberia. Soil Sci. Plant Nutrition, 60, 242-253. https://doi. org/10.1080/00380768.2014.883486
- Furukawa Y., Inubushi K., Ali M., Itang A.M., and Tsuruta H., 2005. Effect of changing groundwater levels caused by land-use changes on greenhouse gas fluxes from tropical peat lands. Nutrient Cycling in Agroecosystems, 71, 81-91. https://doi.org/10.1007/s10705-004-5286-5
- Hadi A., Inubushi K., Furukawa Y., Purnomo E., Rasmadi M., and Tsuruta H., 2005. Greenhouse gas emissions from tropical peatlands of Kalimantan, Indonesia. Nutrient Cycling in Agroecosystems, 71, 73-80. https://doi. org/10.1007/s10705-004-0380-2
- Hashidoko Y., Takakai F., Toma Y., Darung U., Melling L., Tahara S., and Hatano R., 2008. Emergence and behaviors of acid-tolerant *Janthinobacterium* sp. that evolves N₂O from deforested tropical peatland. Soil Biol. Biochemistry, 40, 116-125. https://doi.org/10.1016/j. soilbio.2007.07.014
- IPCC, 2013. Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change (Eds T.F. Stocker, D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley). Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA. https://doi.org/10.1017/ cbo9781107415324.023
- IPCC, 2014. Climate Change 2014: Synthesis Report. Contribution of Working Groups I, II and III to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change (Core Writing Team, Eds R.K. Pachauri and L.A. Meyer). IPCC, Geneva, Switzerland. https://doi.org/10.1017/ cbo9781107415416
- Ishikura K., Darung U., Inoue T., and Hatano R., 2018. Variation in soil properties regulate greenhouse gas fluxes

and global warming potential in three land use types on tropical peat. Atmosphere, 9, 465. https://doi.org/10.3390/atmos9120465

- Ishikura K., Yamada H., Toma Y., Takakai F., Morishita T., Darung U., Limin A., Limin S.H., and Hatano R., 2017. Effect of groundwater level fluctuation on soil respiration rate of tropical peatland in Central Kalimantan, Indonesia. Soil Sci. Plant Nutrition, 63, 1-13. https://doi.org/10.1080/0 0380768.2016.1244652
- Jauhiainen J., Limin S., Silvennoinen H., and Vasander H., 2008. Carbon dioxide and methane fluxes in drained tropical peat before and after hydrological restoration. Ecology, 89, 3503-3514. https://doi.org/10.1890/07-2038.1
- Kelly T.J., Baird A.J., Roucoux K.H., Baker T.R., Honorio C.E.N., Lawson I.T., and Ríos M., 2014. The high hydraulic conductivity of three wooded tropical peat swamps in northeast Peru: Measurements and implications for hydrological function. Hydrological Processes, 28, 3373-3387. https://doi.org/10.1002/hyp.9884
- Kolb S. and Horn M.A., 2012. Microbial CH₄ and N₂O consumption in acidic wetlands. Frontiers in Microbiology, 3, 1-8. https://doi.org/10.3389/fmicb.2012.00078
- Linn D.M. and Doran J.W., 1984. Effect of water-filled pore space on carbon dioxide and nitrous oxide production in tilled and nontilled soils. Soil Sci. Soc. Am. J., 48, 1267-1272. https://doi.org/10.2136/sssaj1984.03615995004800060013x
- Lloyd J. and Taylor J.A., 1994. On the temperature-dependence of soil respiration. Functional Ecology, 8, 315-323.
- McMillan A.M.S., Pal P., Phillips R.L., Palmada T., Berben P.H., Jha N., Saggar S., Luo J., 2016. Can pH amendments in grazed pastures help reduce N₂O emissions from denitrification? – The effects of liming and urine addition on the completion of denitrification in fluvial and volcanic soils. Soil Biology and Biochemistry, 93, 90-104. https:// doi.org/10.1016/j.soilbio.2015.10.013
- Melling L., 2016. Peatland in Malaysia. Tropical Peatland Ecosystems, 59-73. https://doi.org/10.1007/978-4-431-55681-7 4
- Melling L., Hatano R., and Goh K.J., 2005. Soil respiration from three ecosystems in tropical peatland of Sarawak, Malaysia. Tellus, 57, 1-11. https://doi.org/10.1111/j.1600-0889.2005.00129.x
- Melling L., Hatano R., and Goh K.J., 2007. Nitrous oxide emissions from three ecosystems in tropical peatland of Sarawak, Malaysia. Soil Sci. Plant Nutrition, 53: 792-805. https://doi. org/10.1111/j.1747-0765.2007.00196.x
- Mu Z.. Huang A., Kimura S.D., Jin T., Wei S., and Hatano R., 2009. Linking N₂O emission to soil mineral N as estimated by CO₂ emission and soil C/N ratio. Soil Biol. Biochem., 41, 2593-2597. https://doi.org/10.1016/j.soilbio.2009.09.013
- Mu Z., Huang A., Ni J., and Xie D., 2014. Linking annual N₂O emission in organic soils to mineral nitrogen input as estimated by heterotrophic respiration and soil C/N ratio. Plos One, 9, e96572. https://doi.org/10.1371/journal.pone. 0096572
- Mukumbuta I., Uchida Y., and Hatano R., 2018. Evaluating the effect of liming on N₂O fluxes from denitrification in an

Andosol using the acetylene inhibition and N-15 isotope tracer methods. Biol. Fertility Soils, 54, 71-81. https://doi. org/10.1007/s00374-017-1239-4

- Page S.E. and Baird A.J., 2016. Peatlands and global change: response and resilience. Annual Review of Environment and Resources, 41, 35-57. https://doi.org/10.1146/annurevenviron-110615-085520
- Pangala S.R., Moore S., Hornibrook E.R., Gauci V., 2013. Trees are major conduits for methane egress from tropical forested wetlands. New Phytologist, 197, 524-31. https:// doi.org/10.1111/nph.12031
- Repo M.E., Susiluoto S., Lind S.E., Jokinen S., Elsakov V., Biasi C., Virtanen T., and Martikainen P.J., 2009. Large N₂O emissions from cryoturbated peat soil in tundra. Nature Geoscience, 2, 189-192. https://doi.org/10.1038/ngeo434
- Saggar S., Jha N., Deslippe J., Bolan N.S., Luo J., Giltrap D.L., Kim D.G., Zaman M., and Tillman R.W., 2013. Denitrification and N₂O:N₂ production in temperate grasslands: Processes, measurements, modelling and mitigating negative impacts. Sci. Total Environ., 465, 173-195. https:// doi.org/10.1016/j.scitotenv.2012.11.050
- Schütz H., Seiler W., and Conrad R., 1989. Processes involved in formation and emission of methane in rice paddies. Biogeochemistry, 7, 33-53. https://doi.org/10.1007/bf00000896
- Šimek M., Jíšová L., and Hopkins D.W., 2002. What is the so-called optimum pH for denitrification in soil? Soil Biol. Biochem., 34,1227-1234. https://doi.org/10.1016/s0038-0717(02)00059-7
- Spahni R., Wania R., Neef L., van Weele M., Pison I., Bousquet P., Frankenberg C., Foster P.N., Joos F., Prentice I.C., and van Velthoven P., 2011. Constraining global methane emissions and uptake by ecosystems, Biogeosciences, 8, 1643-1665. https://doi.org/10.5194/bg-8-1643-2011
- Takahashi H., 1999. Hydrological and meteorological environments of inland peat swamp forest in central Kalimantan, Indonesia with special reference to the effects of forest fire. Tropics, 9(1), 17-25. https://doi.org/10.3759/tropics.9.17
- Takakai F., Morishita T., Hashidoko Y., Darung U., Kuramochi K., Dohong S., Limin S.H., and Hatano R., 2006. Effects of agricultural land-use change and forest fire on N₂O emission from tropical peatlands, Central Kalimantan, Indonesia. Soil Sci. Plant Nutrition, 52, 662-674. https://doi. org/10.1111/j.1747-0765.2006.00084.x
- Takai Y., 1970. The mechanism of methane fermentation in flooded paddy soil. Soil Sci. Plant Nutrition, 16, 238-244. https:// doi.org/10.1080/00380768.1970.10433371
- Toma Y., Takakai F., Darung U., Kuramochi K., Limin S.H., Dohong S., and Hatano R., 2011. Nitrous oxide emission derived from soil organic matter decomposition from tropical agricultural peat soil in central Kalimantan, Indonesia. Soil Sci. Plant Nutrition, 57, 436-451. https://doi.org/10.10 80/00380768.2011.587203
- UNEP, **2013.** Drawing Down N₂O To Protect Climate and the Ozone Layer. United Nations Environment Programme (UNEP), Nairobi, Kenya.
- van der Werf G.R., Morton D.C. DeFries R.S. Olivier J.G.J., Kasibhatla P.S., Jackson R.B., Collatz G.J. and

Randerson J.T., 2009. CO₂ emissions from forest loss. Nature Geoscience, 2, 737-738. https://doi.org/10.1038/ ngeo671

Yanai Y., Toyota K., Morishita T., Takakai F., Hatano R., Limin S.H., Darung U., and Dohong S., 2007. Fungal N₂O production in an arable peat soil in Central Kalimantan, Indonesia. Soil Sci. Plant Nutrition, 53, 806-811. https:// doi.org/10.1111/j.1747-0765.2007.00201.x

Yu Z., Loise J., Brosseau D.P., Beilman D.W. and Hunt S.J.,

2010. Global peatland dynamics since the Last Glacial Maximum. Geophysical Res. Letters, 37, L13402. https://doi.org/10.1029/2010gl043584

Zhou L., Zhou X., Zhang B., Lui M., Luo Y., Liu L., and Li B.,
2014. Different responses of soil respiration and its components to nitrogen addition among biomes: A meta-analysis.
Global Change Biol., 20, 2332-2343. https://doi. org/10.1111/gcb.12490